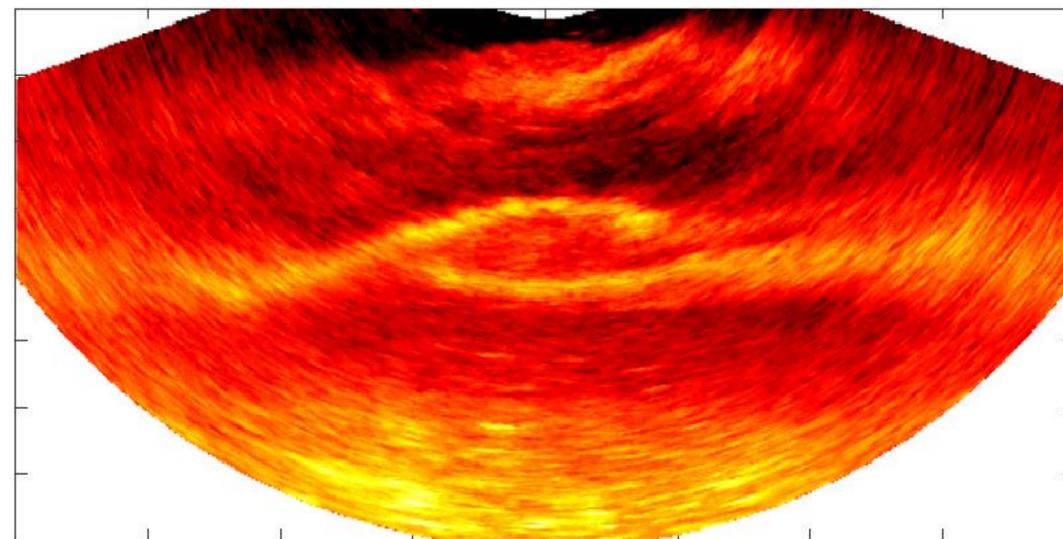


# The structure and lifecycle of stratified mixing by shear instability in continuously forced shear flows

strongly-stratified estuary



multibeam sonar



Signal: diapycnal mixing rate of salinity  $\chi_s$

Adrien Lefauve

Chris Bassett

Dan Plotnick

Andone Lavery

Rocky Geyer

IMPERIAL

UNIVERSITY of WASHINGTON

PennState

Oregon State University

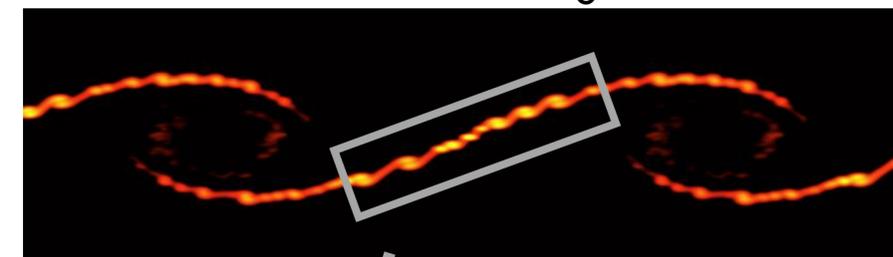
WOODS HOLE OCEANOGRAPHIC INSTITUTION



Natural Environment Research Council

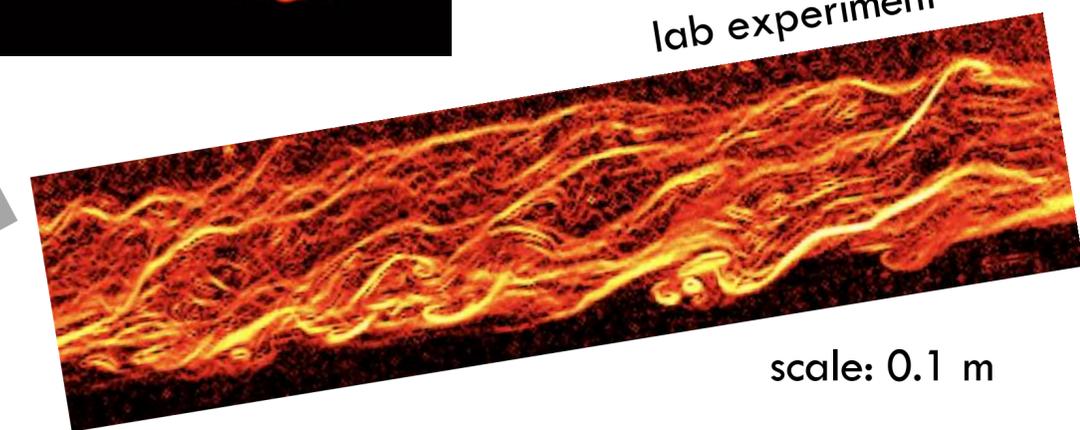


numerical modelling



scale: 1 m

lab experiment



scale: 0.1 m

# The site, transects and instruments

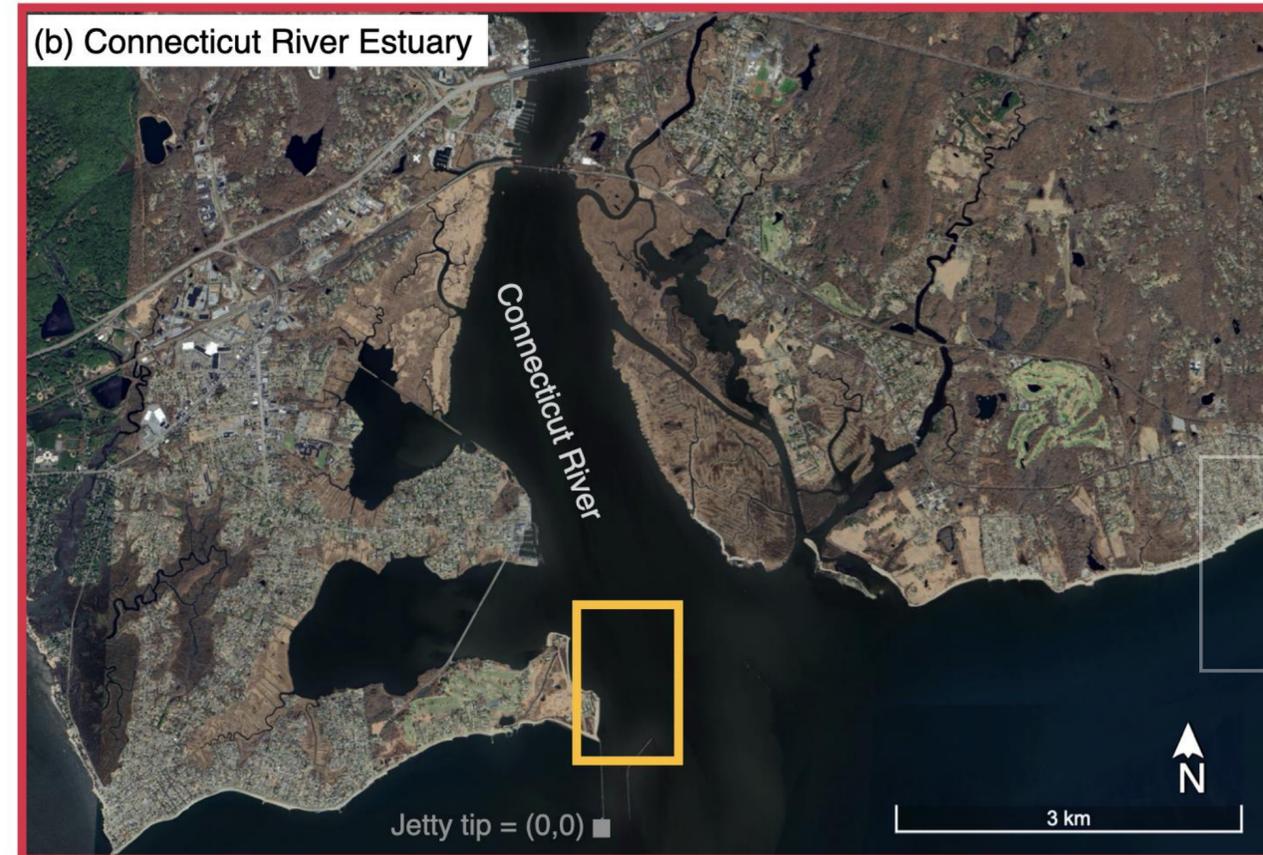
Fieldwork by Rocky Geyer & Andone Lavery (June 2017)



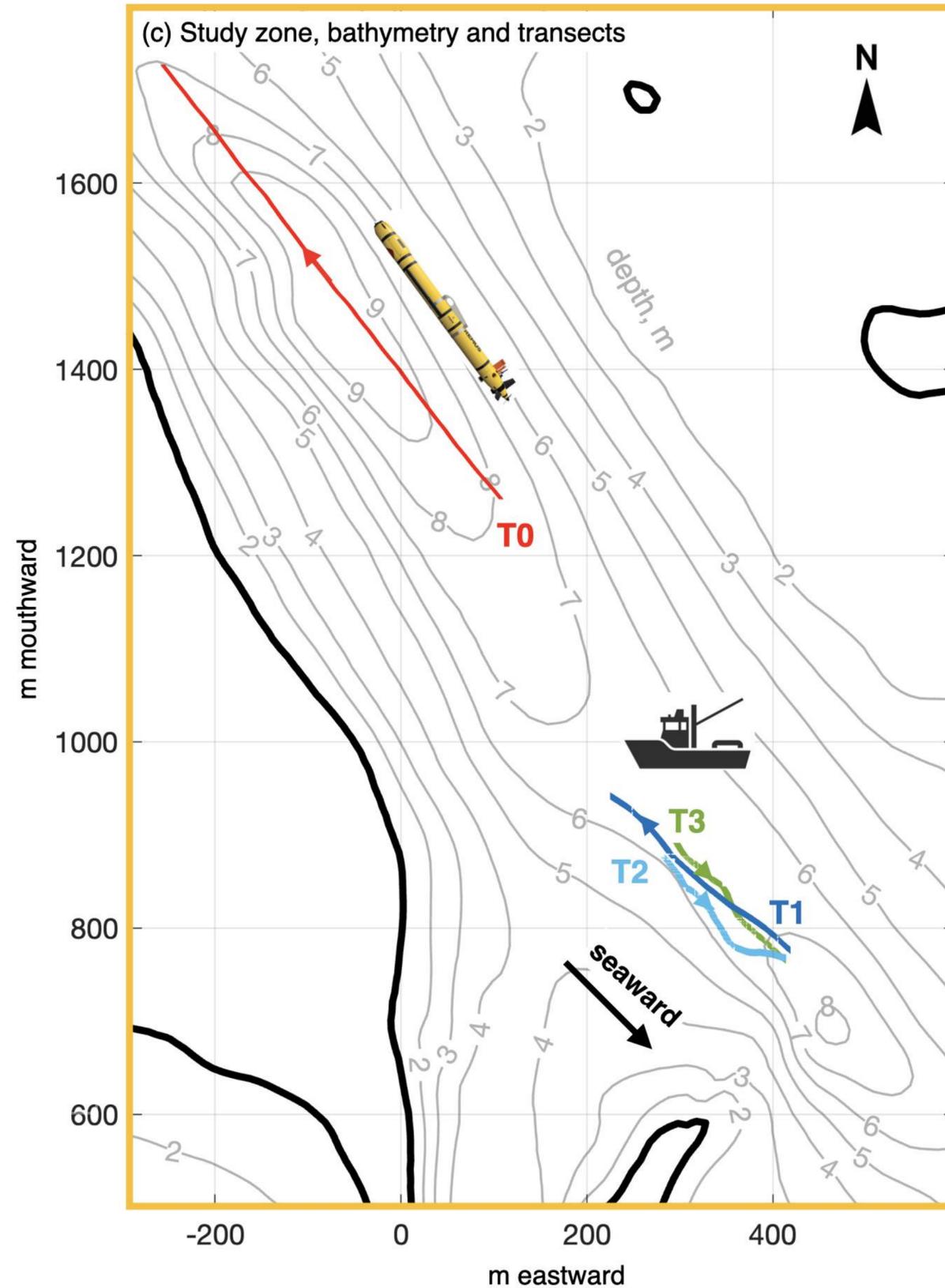
(a) Overview



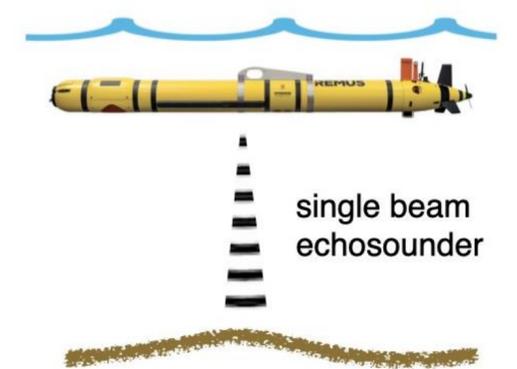
(b) Connecticut River Estuary



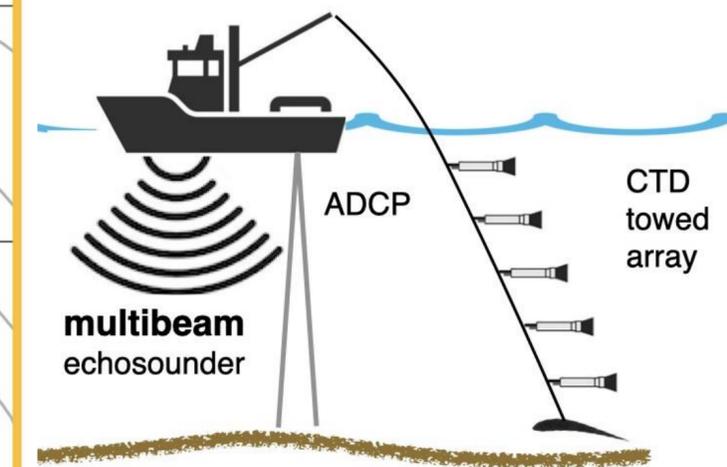
(c) Study zone, bathymetry and transects



REMUS AUV → Transect T0

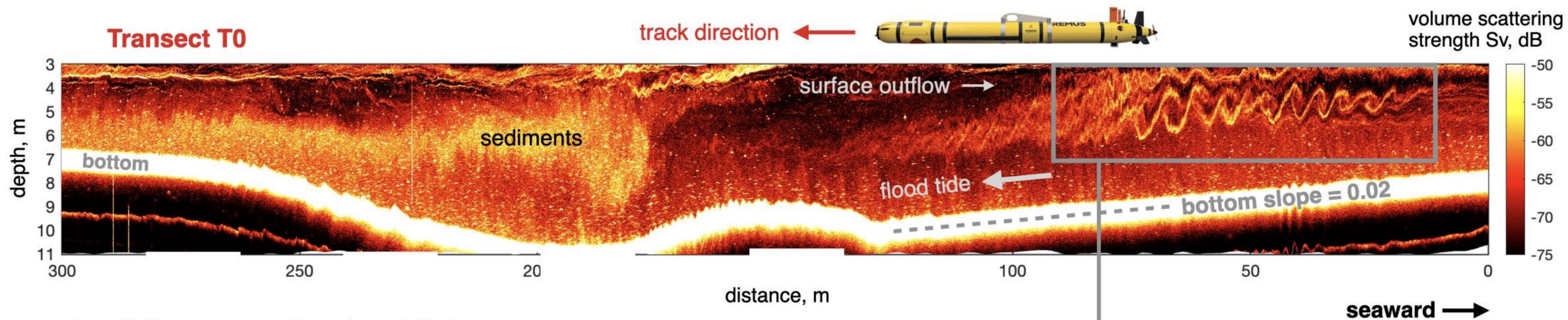


Shipboard → Transects T1 T2 T3

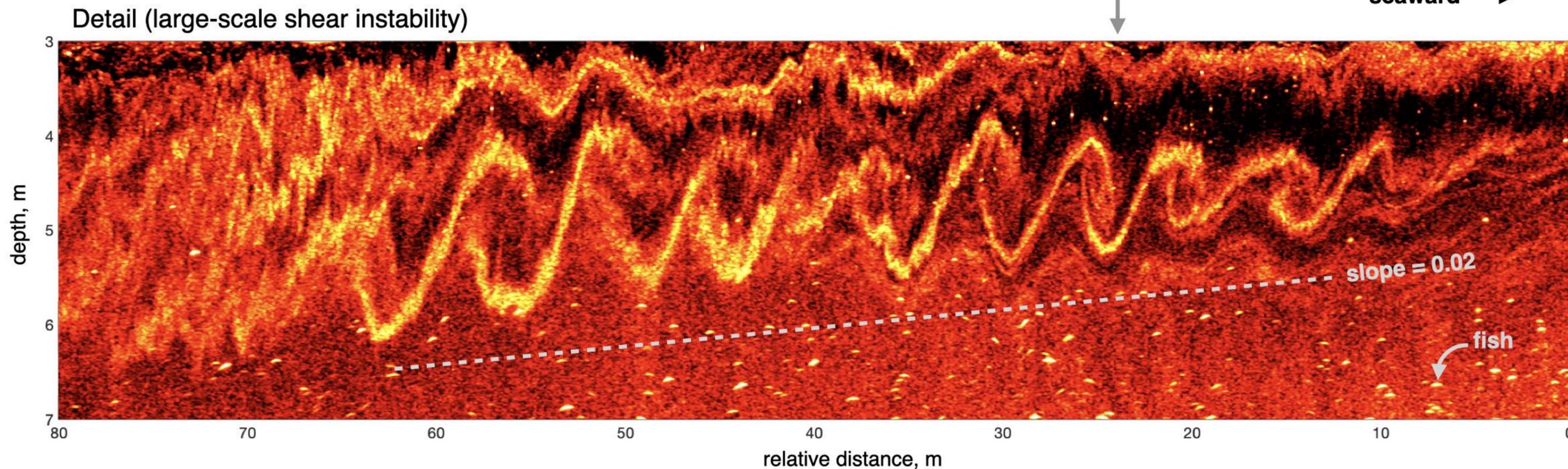


# Instability train at the pycnocline revealed by echosounder

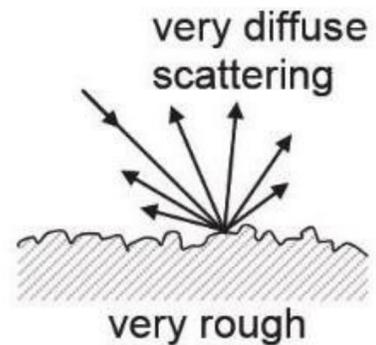
Remus  
Autonomous  
Underwater  
Vehicle  
(AUV)



Vertical  
exaggeration = 5



Principle:  
Bragg scattering



Acoustic backscatter is a proxy for mixing rate

$$S_V = 5 \log_{10} \chi_s + \text{const.}$$

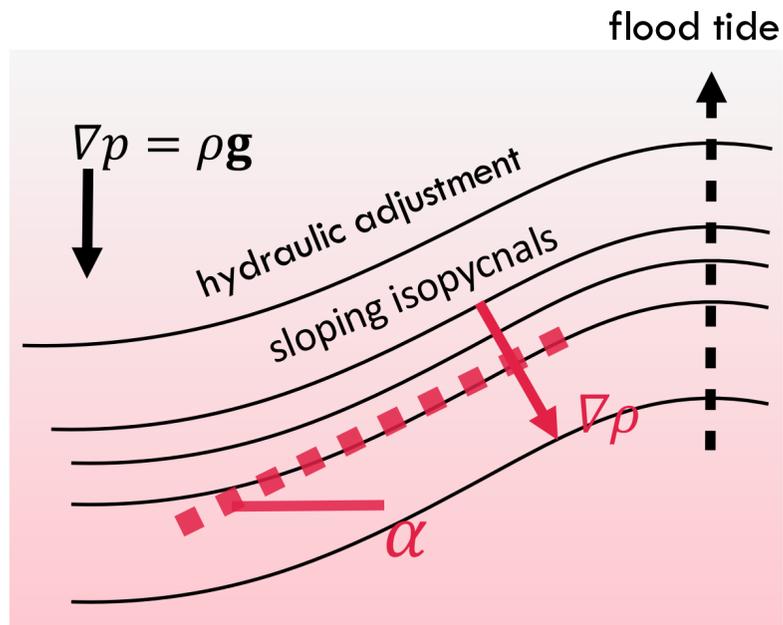
Lavery, Geyer & Scully, J. Acous. Soc. Am. (2013)

$\chi_s = 2D_s |\nabla s'|$  dissipation rate of salinity variance  
~ diapycnal mixing rate

# Continuous shear forcing

Salinity-stratified flow (temperature negligible)

Tidal forcing → **baroclinic shear** generation



$$\frac{D\omega}{Dt} = (\omega \cdot \nabla)\mathbf{u} + \frac{1}{\rho_0^2} \nabla\rho \times \nabla p$$

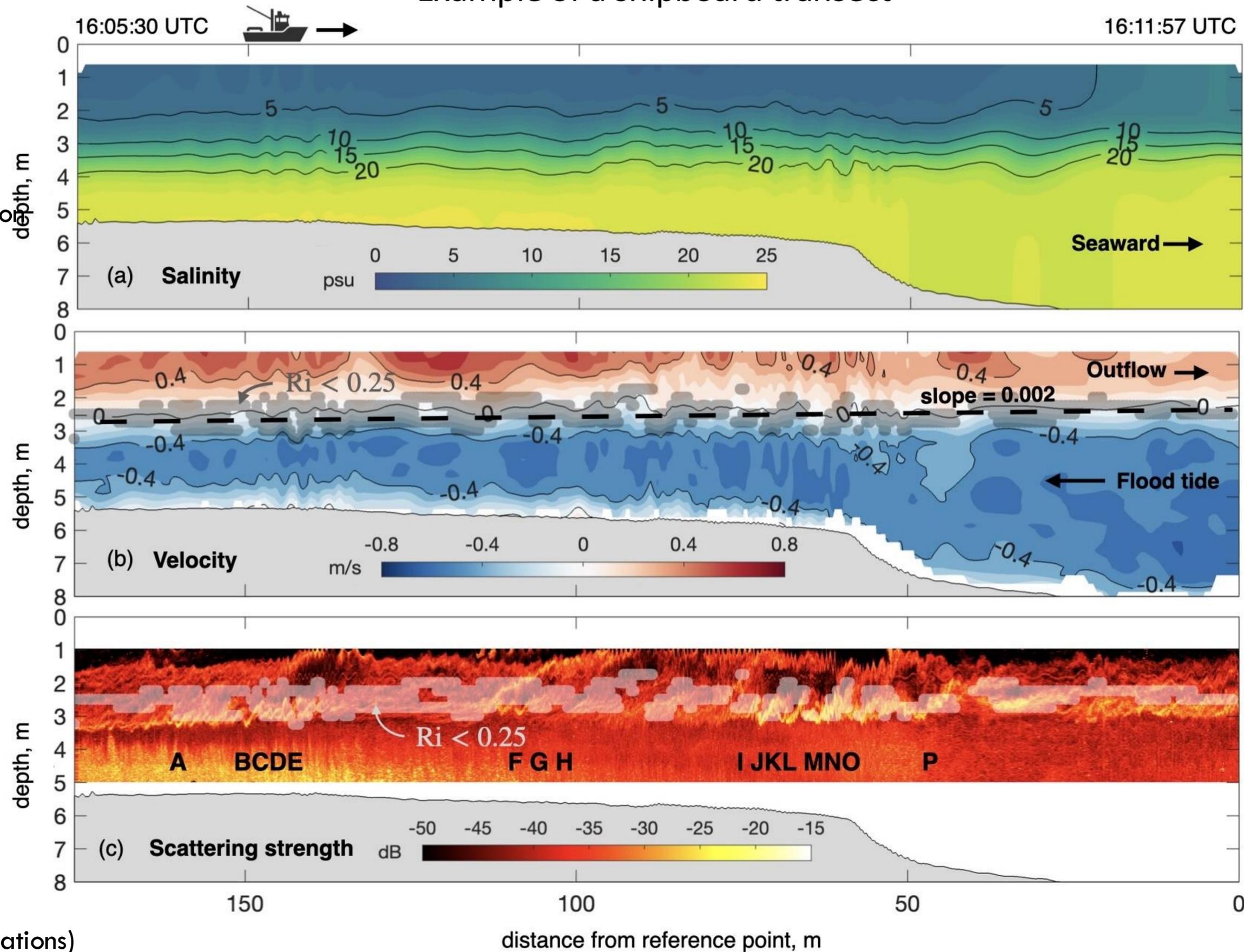
$$\frac{\partial S}{\partial t} \approx N^2 \sin \alpha \approx \alpha N^2$$

$$\text{Drop in Ri} = \frac{N^2}{S^2} < 0.25$$

→ **KH instability** (predictions match observations)

## Transect T2

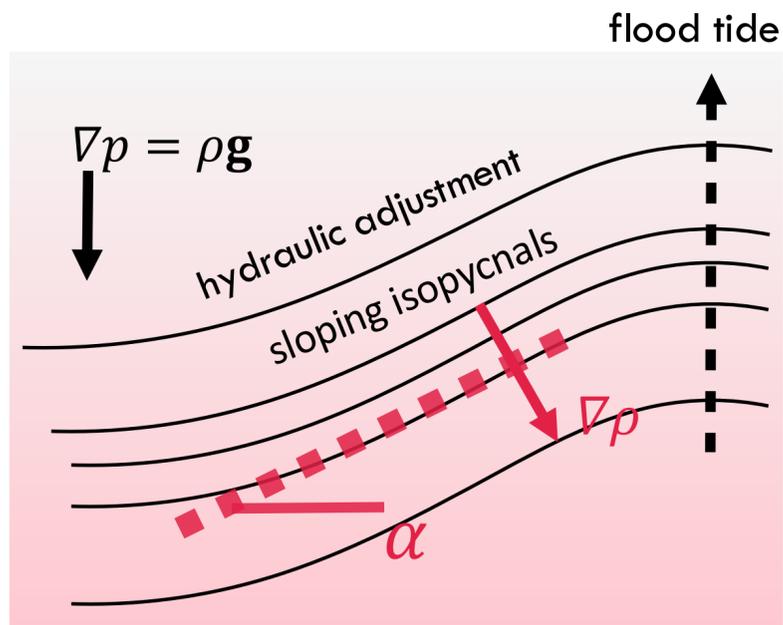
## Example of a shipboard transect



# Continuous shear forcing

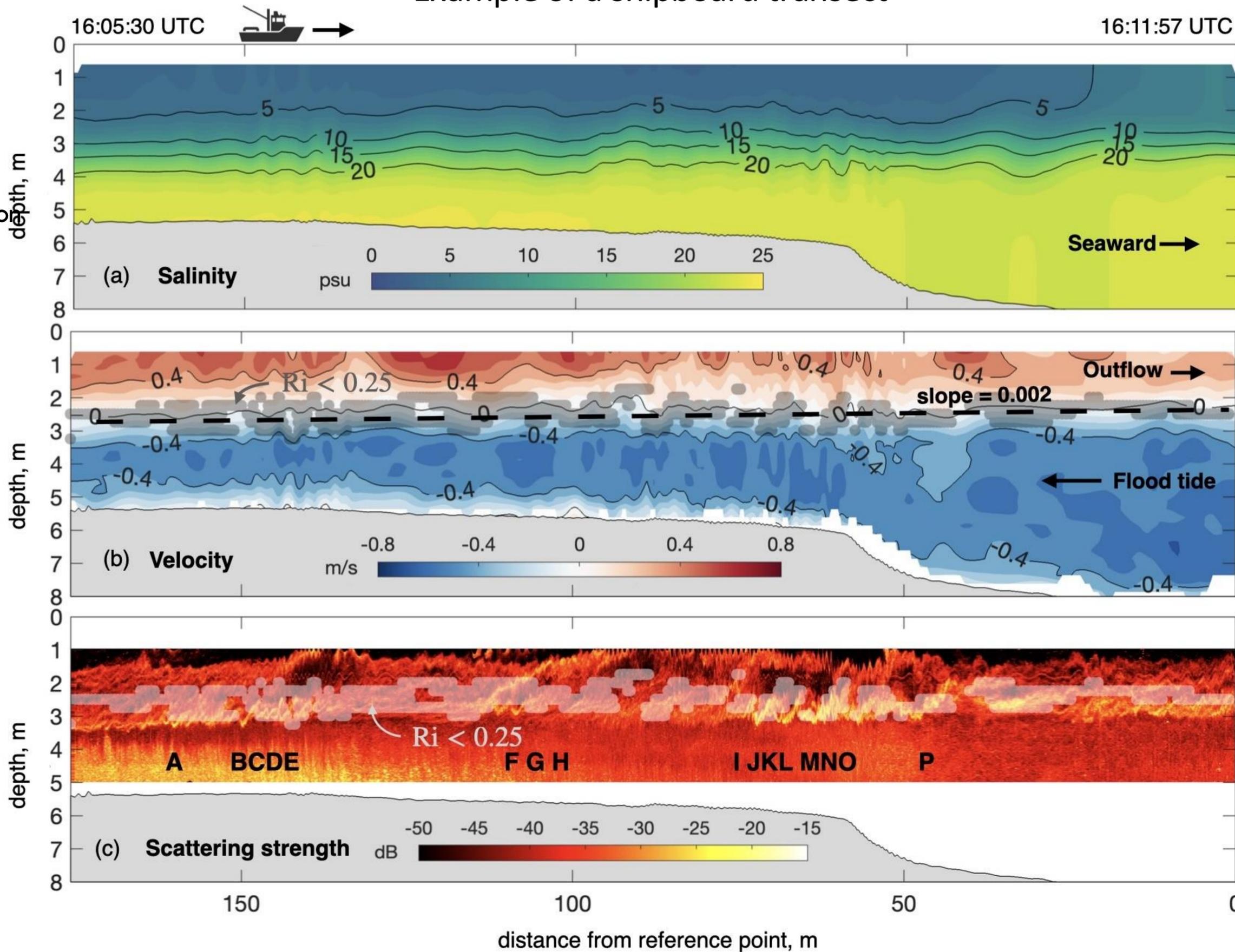
Salinity-stratified flow (temperature negligible)

Tidal forcing → **baroclinic shear** generation

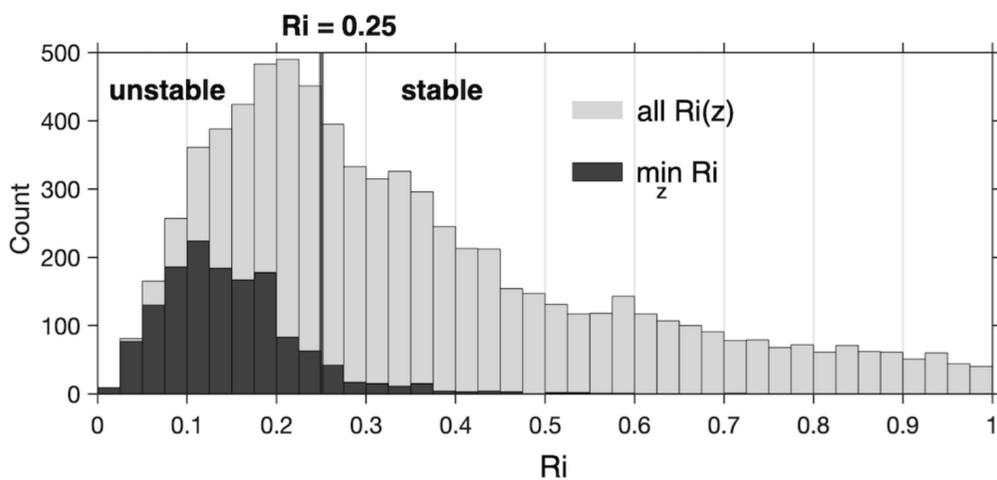


## Transect T2

## Example of a shipboard transect



(b) Ri histogram, depth  $z = 2 - 4$  m



LOW Ri !

# Multibeam reveals the spatial structure and lifecycle of instabilities

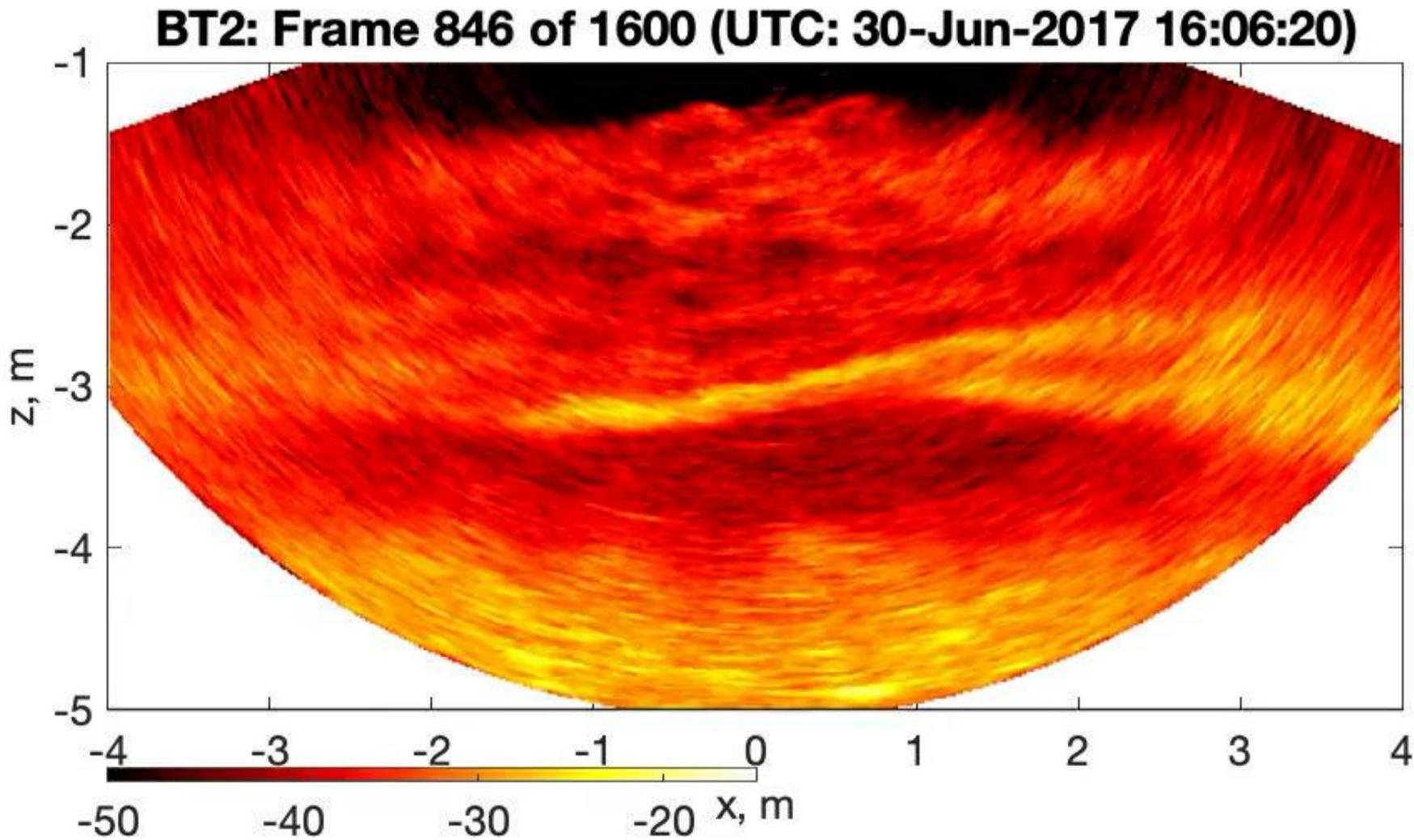
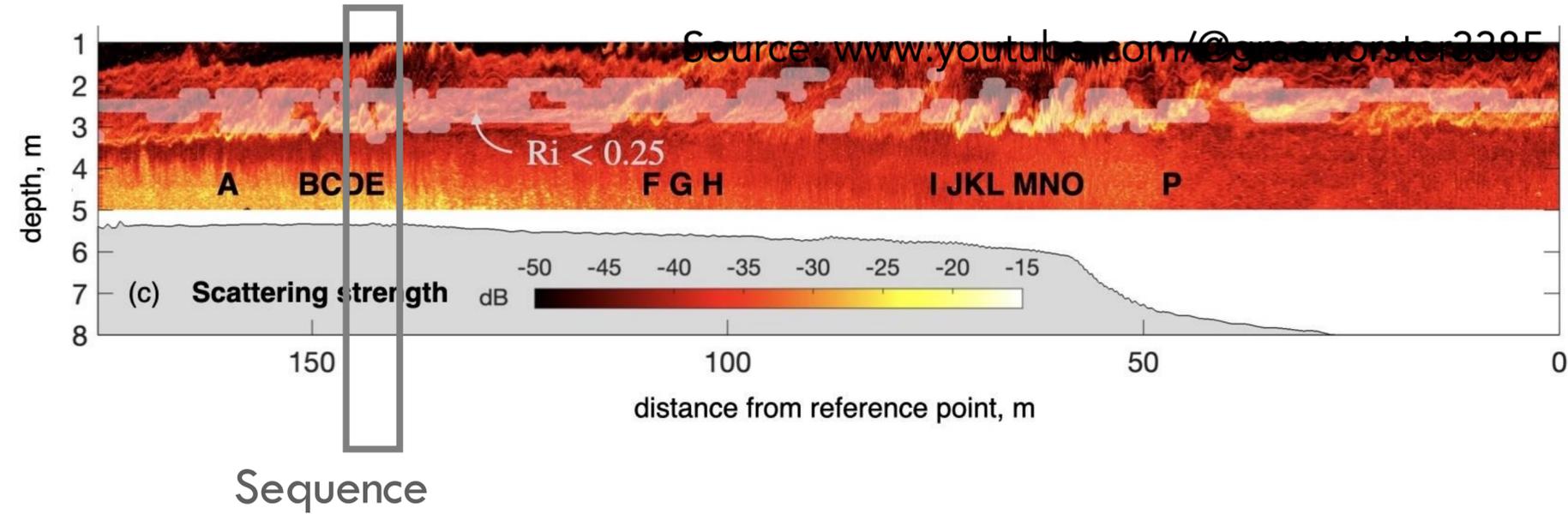
Combine 75 individual beams  $\rightarrow$  120-degree fan

Motion correction (boat heave, pitch and roll)

Multibeam disentangles space and time

1:1 aspect ratio

Real speed



# Multibeam reveals the spatial structure and lifecycle of instabilities

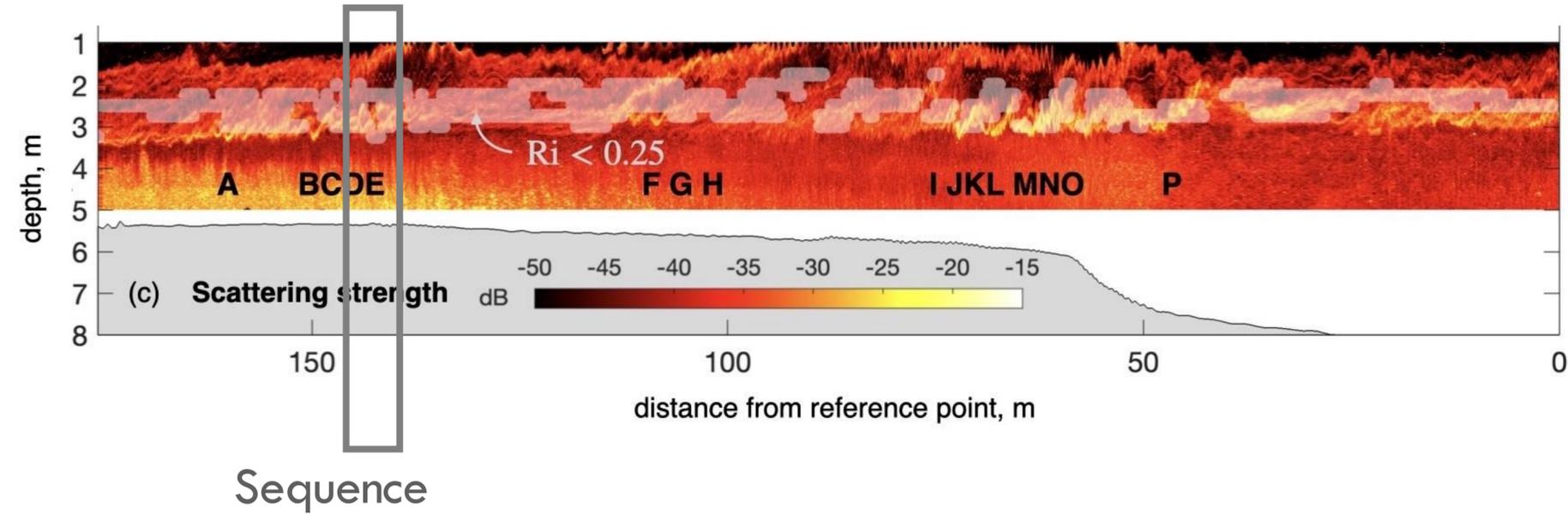
Combine 75 individual beams → 120-degree fan

Motion correction (boat heave, pitch and roll)

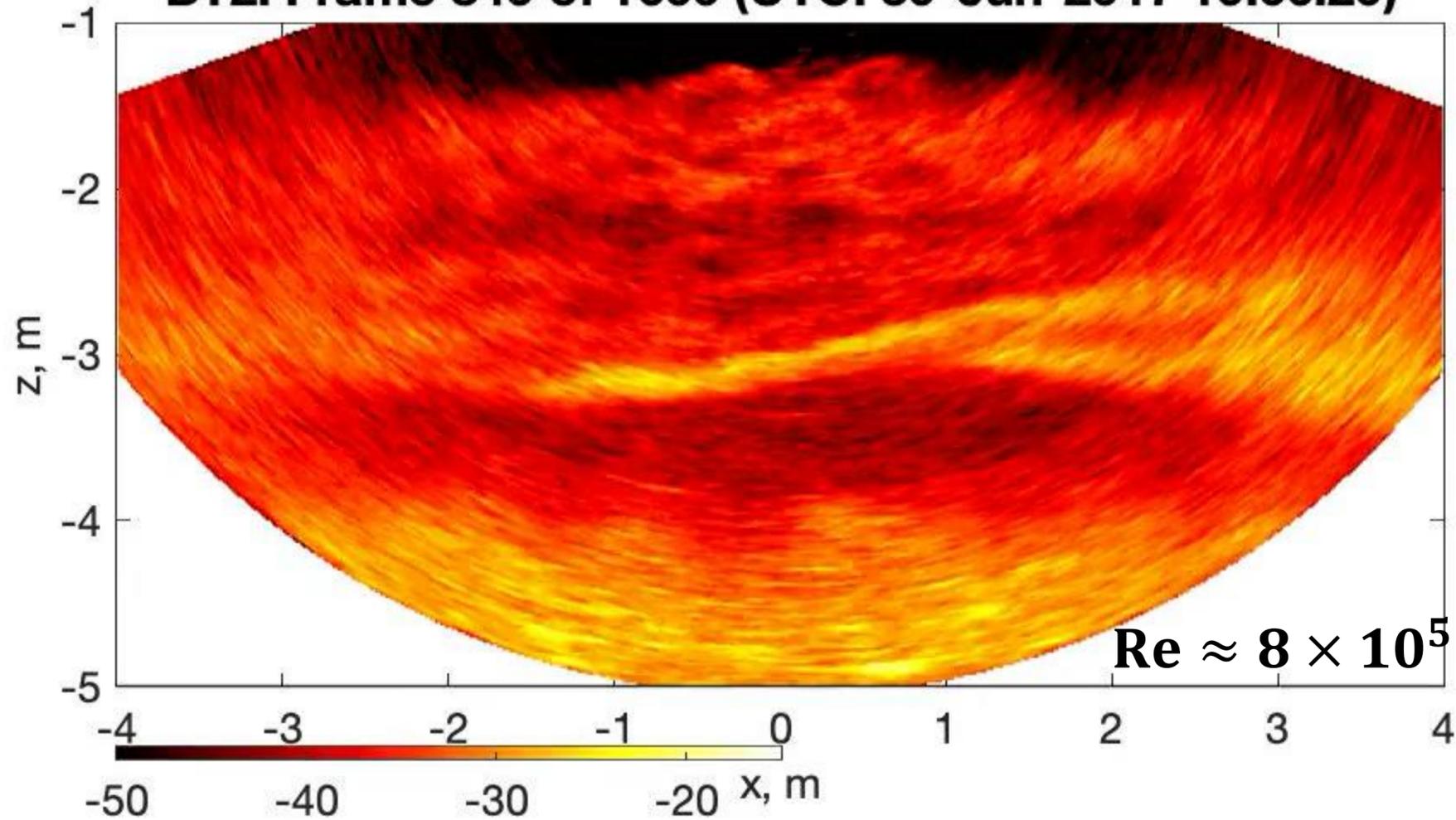
Multibeam disentangles space and time

1:1 aspect ratio

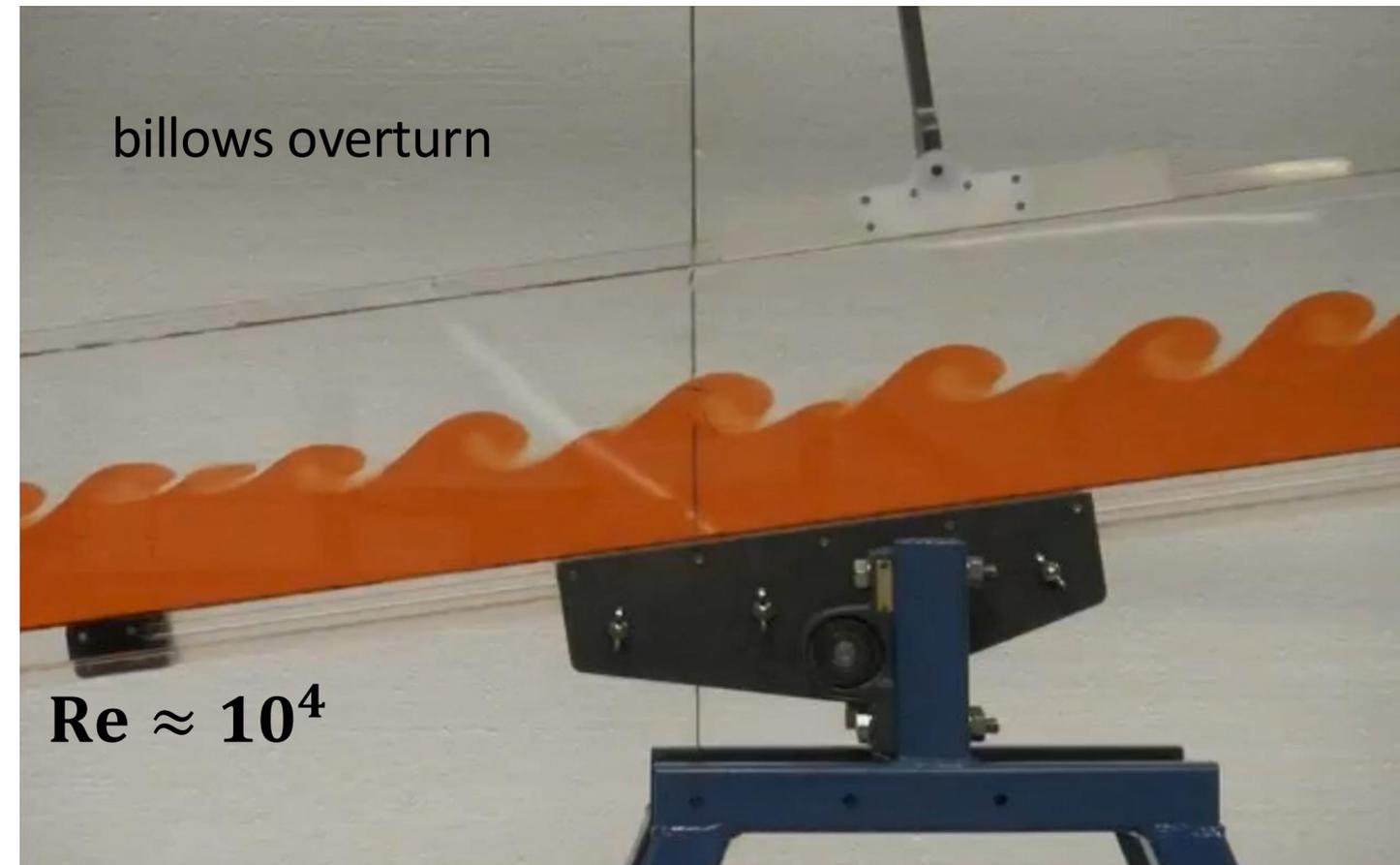
Real speed



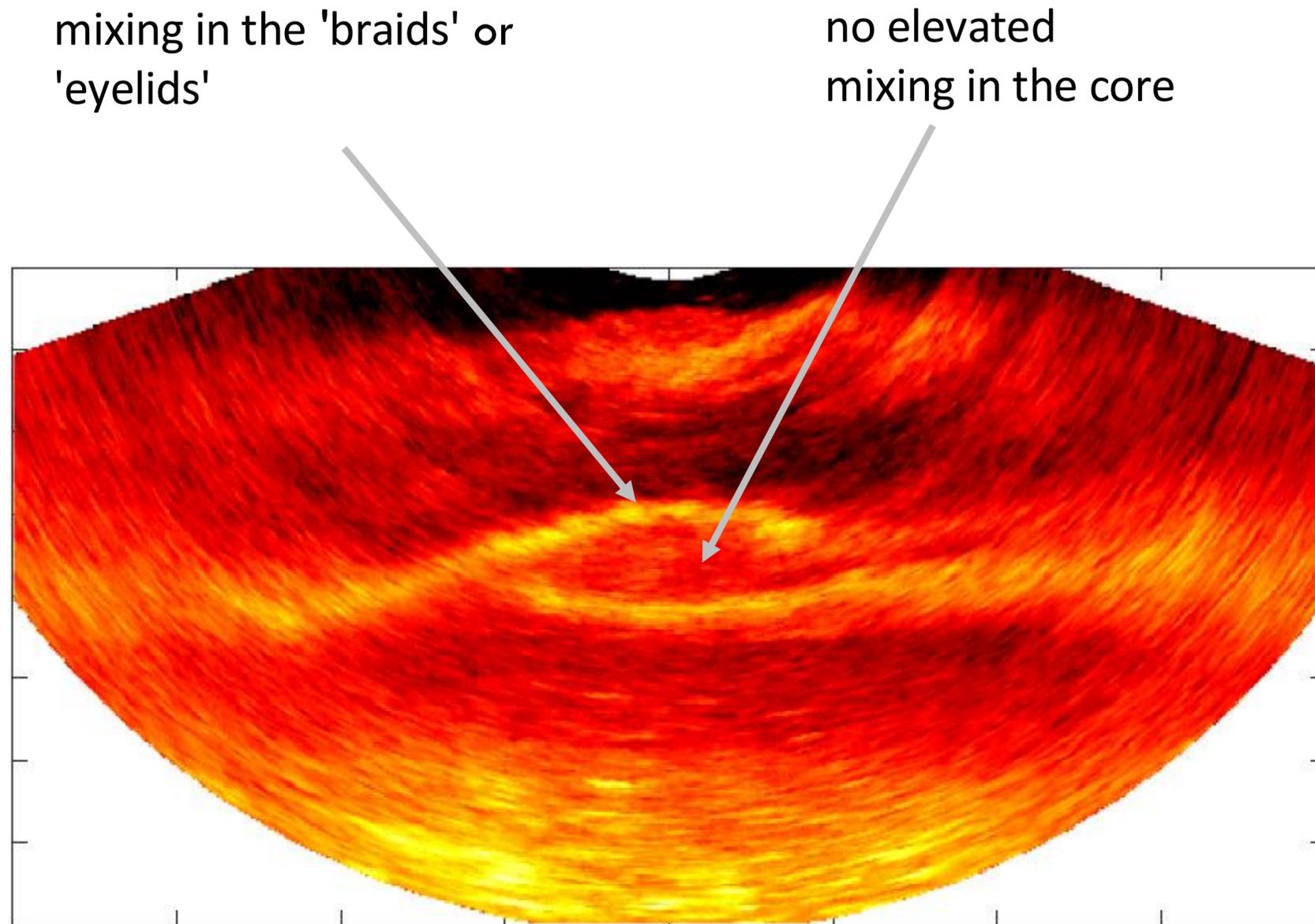
**BT2: Frame 846 of 1600 (UTC: 30-Jun-2017 16:06:20)**



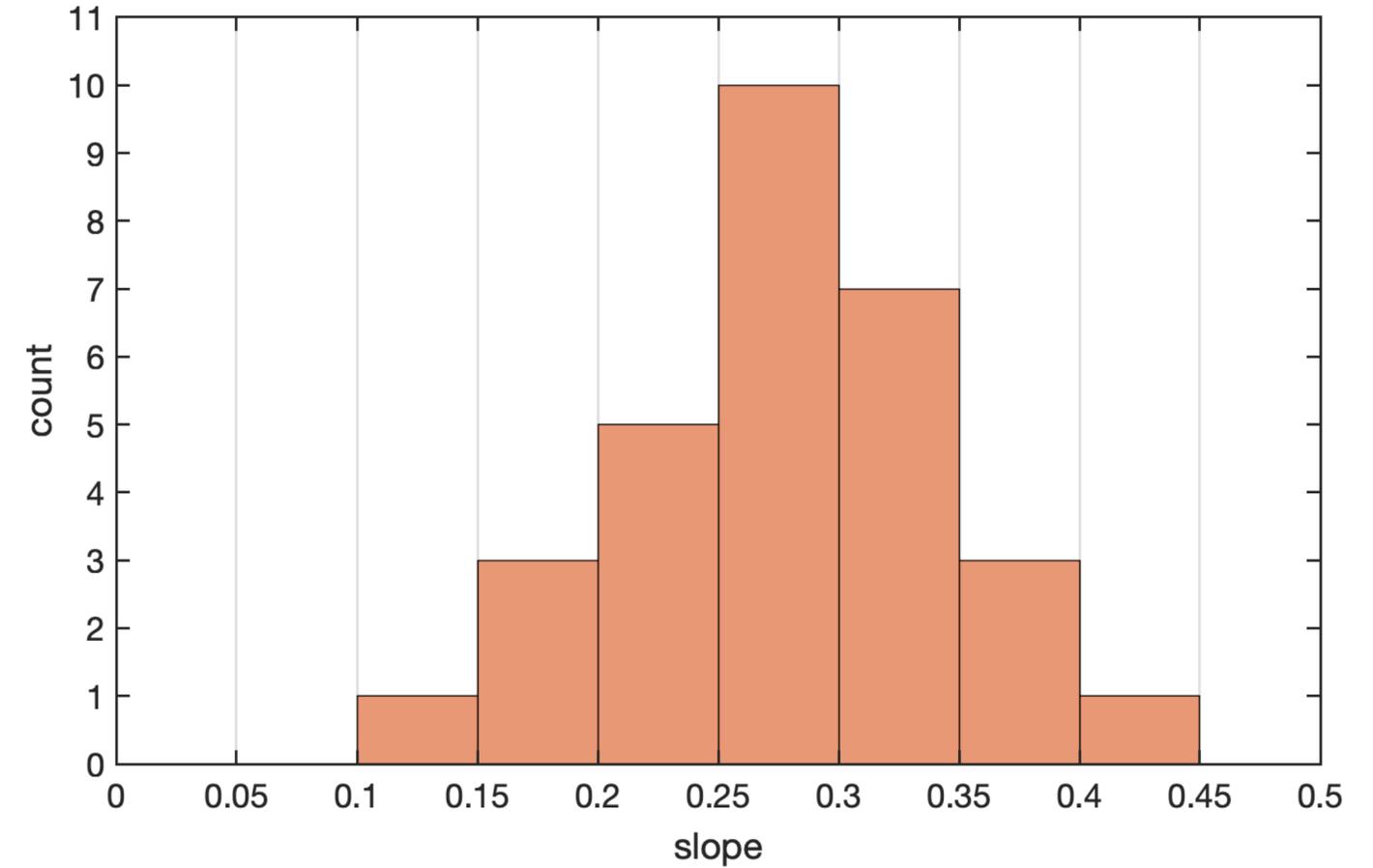
Contrast with KH breaking in the lab



# Turbulent BRAID mixing without CORE overturn



Slopes measured in 30 events from multibeam:



moderate slopes → too shallow to overturn

Why braid turbulence?

Why no core overturn?

# Mechanism: secondary stratified shear instability at high Re

2D simulations

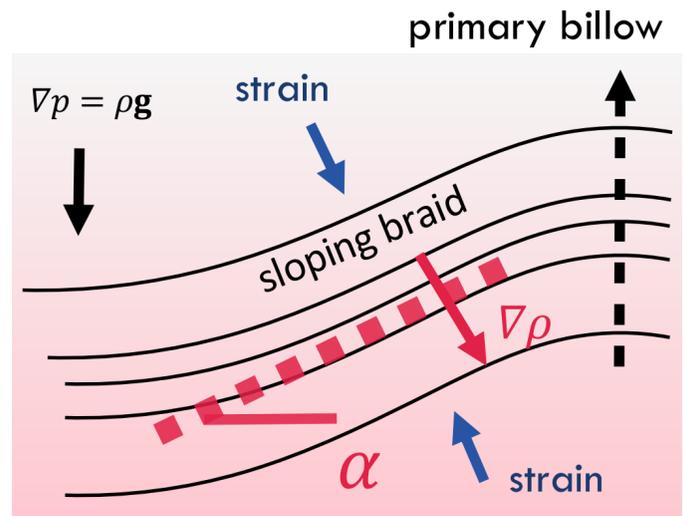
FIELD

$Re = 8 \times 10^5$

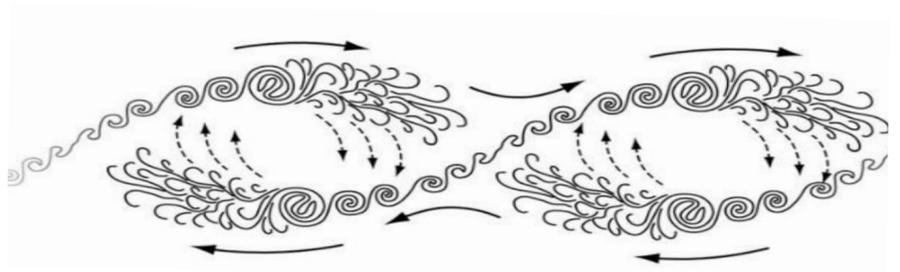
$Ri = 0.15,$

$Pr = 700$

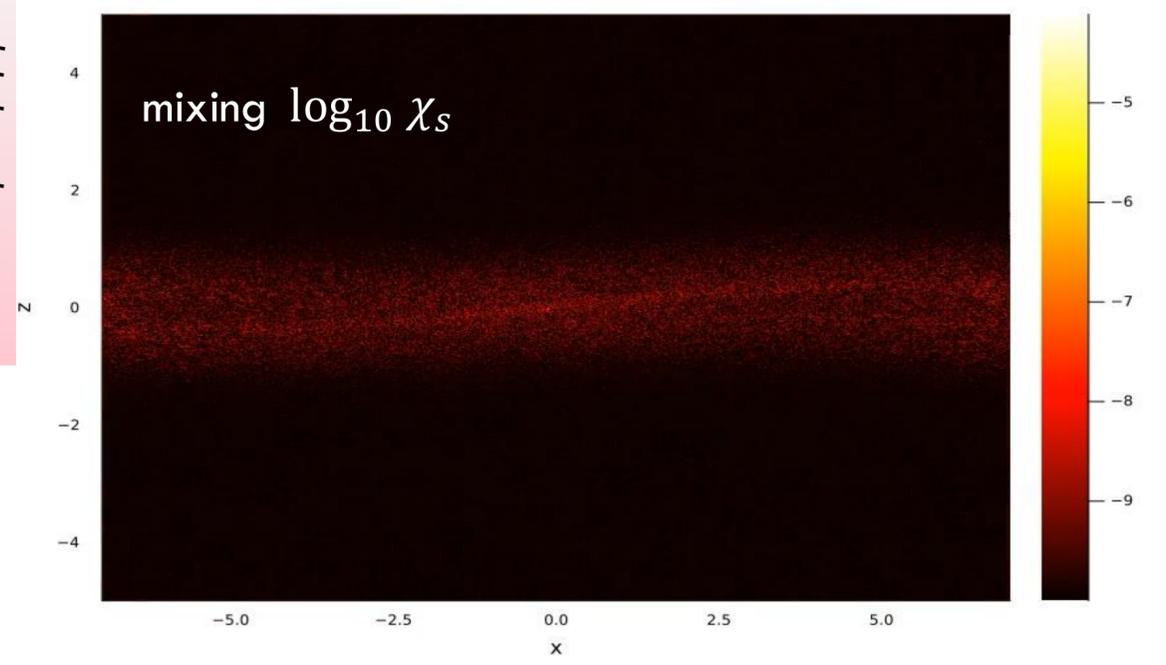
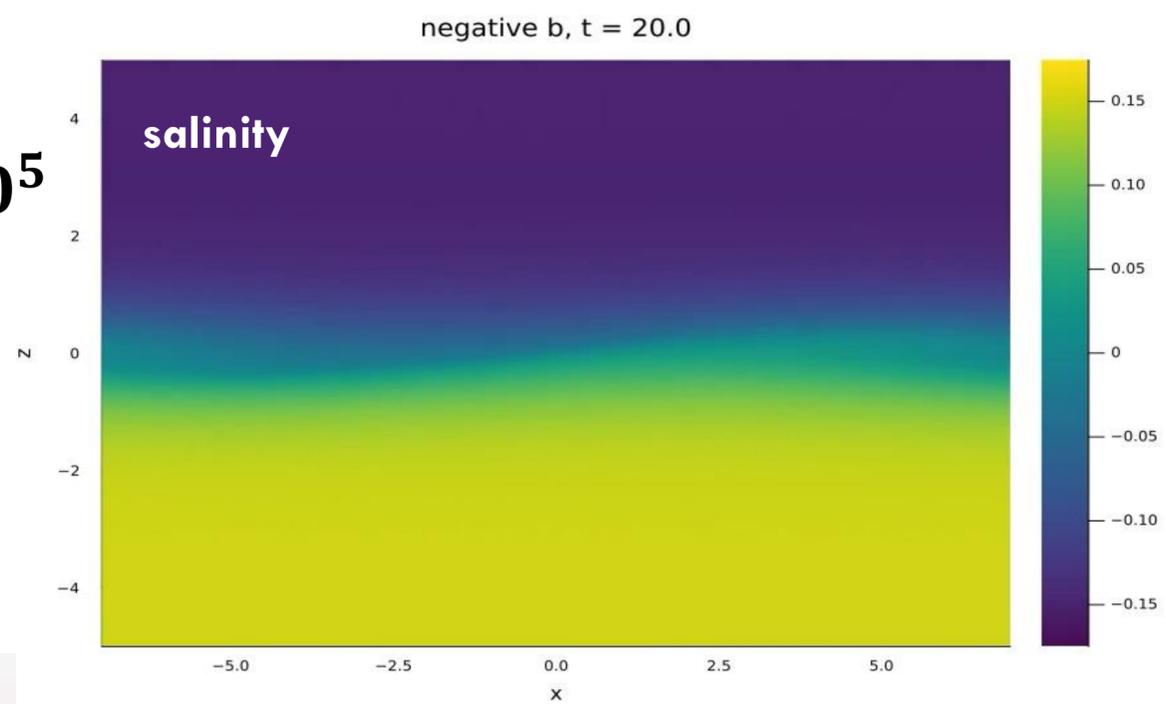
Baroclinic shear intensification:



Corcos & Sherman, J. Fluid Mech. (1976)  
 Dritschel et al. J. Fluid Mech. (1991)  
 Staquet, J. Fluid Mech. (1995)  
 Smyth, J. Fluid Mech. (2003)



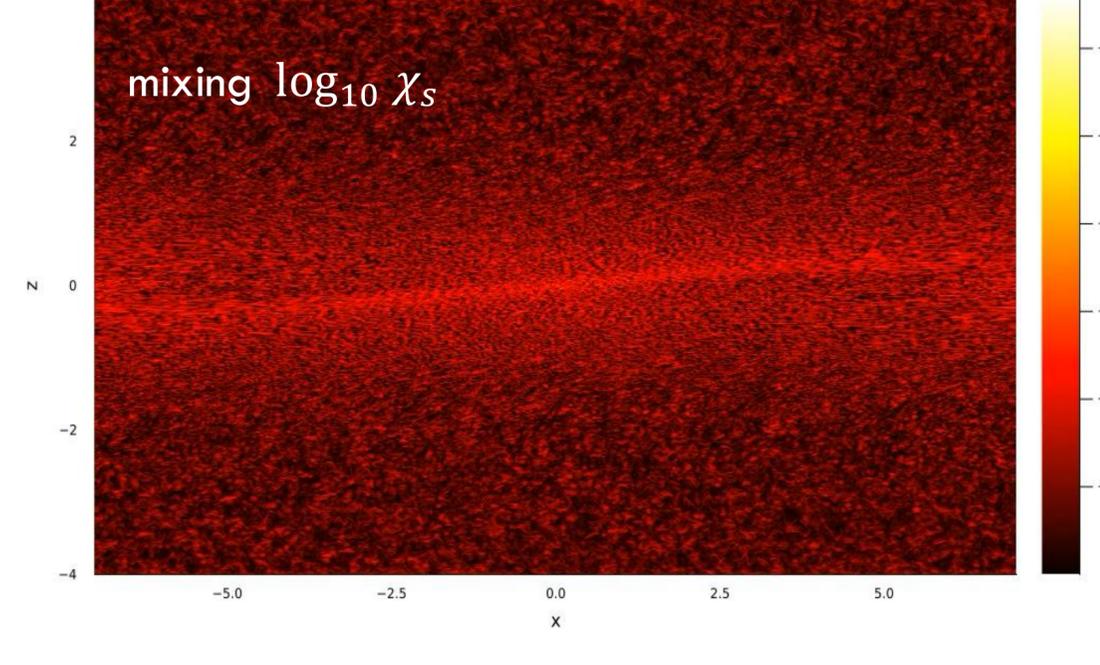
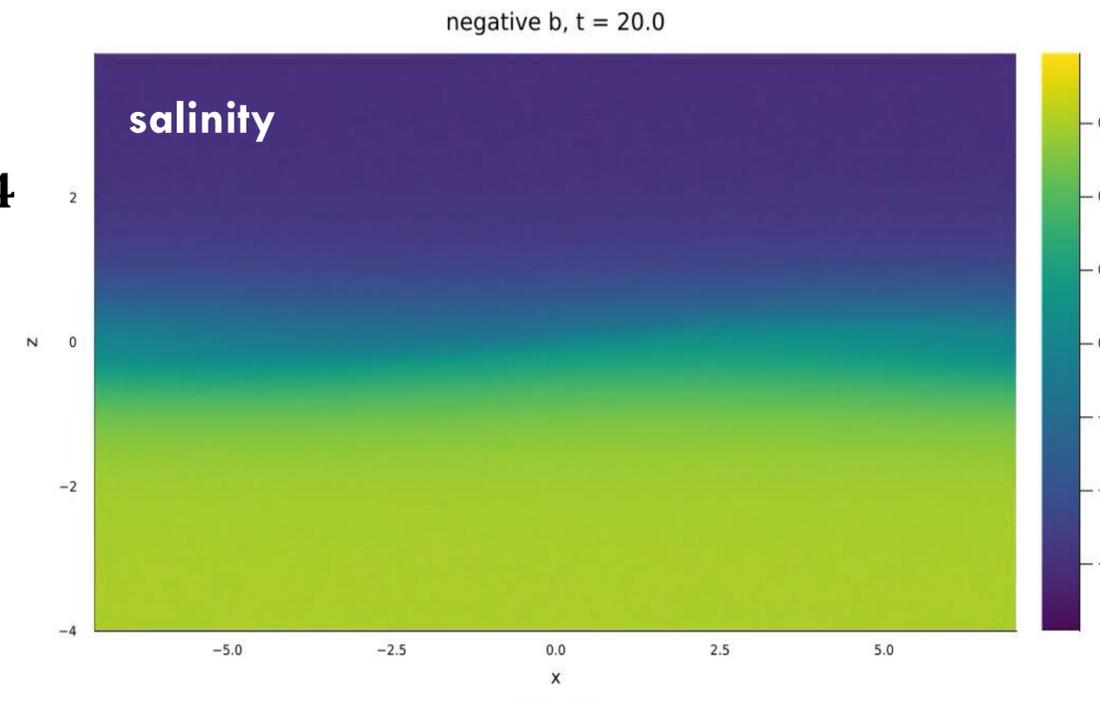
already proposed by  
 Geyer et al. Geophys. Res. Lett. (2010)



LAB

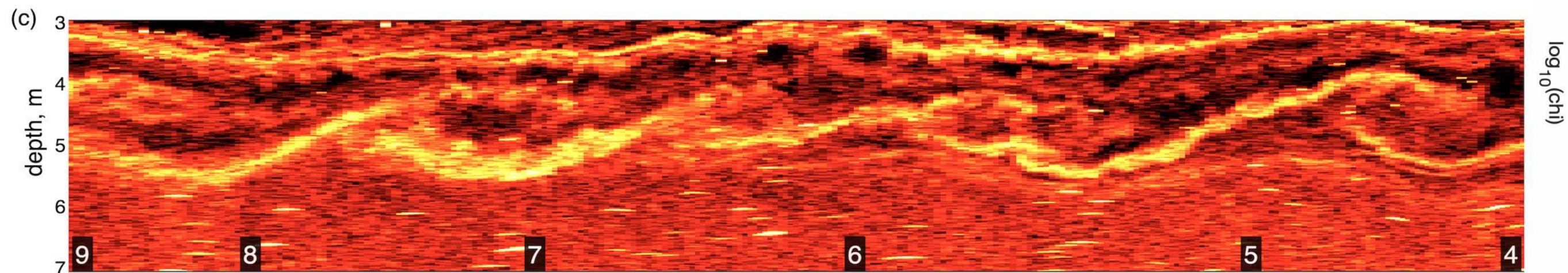
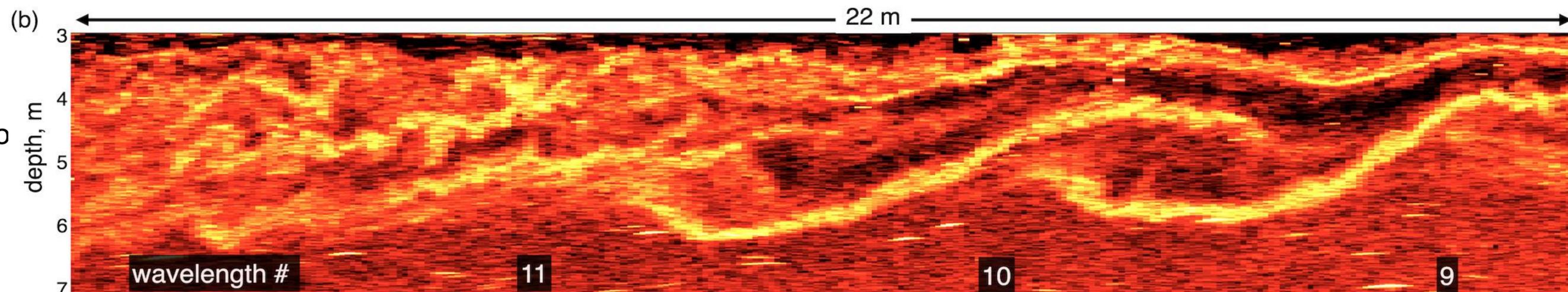
$Re = 10^4$

same Ri, Pr



# Further evidence of braid secondary shear instability

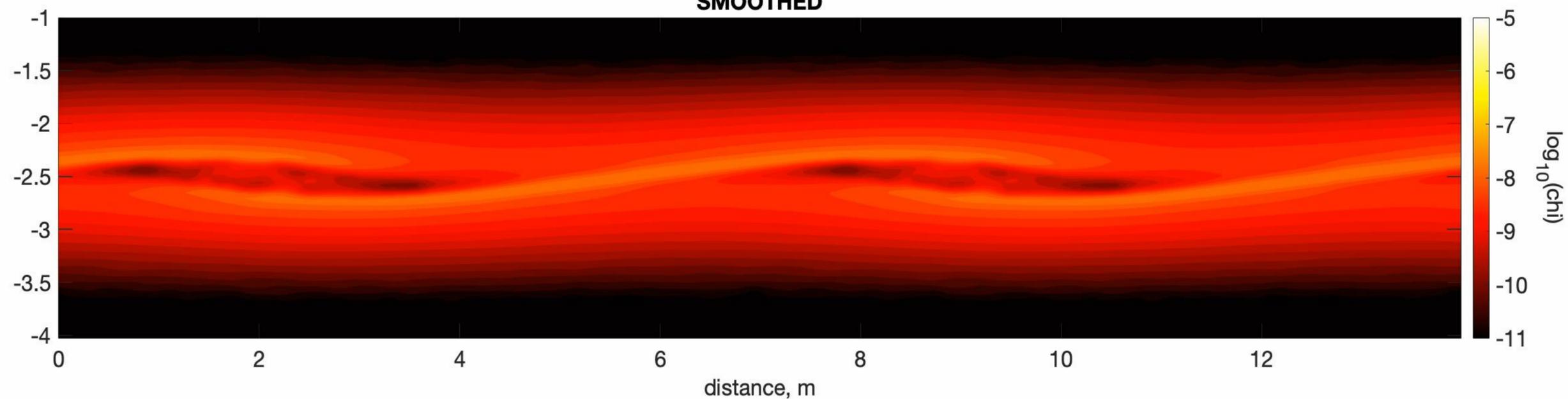
AUV  
echosounder  
1:1 aspect ratio



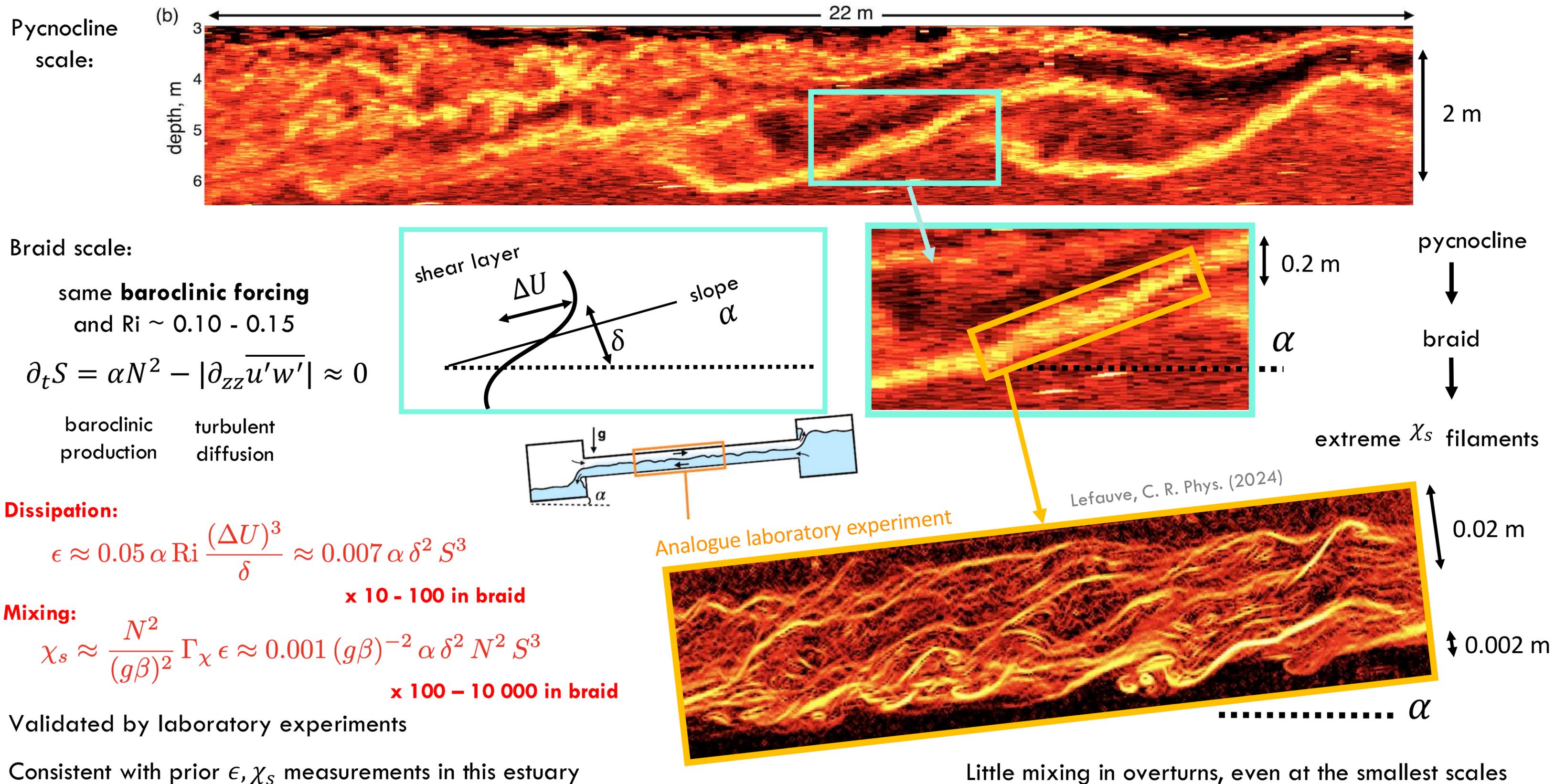
SMOOTHED

2D simulation

smoothed to  
simulate  
acoustic  
backscatter

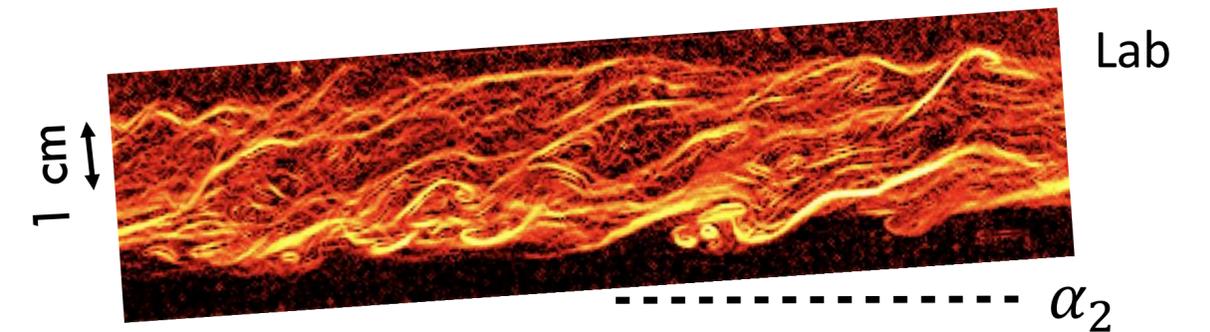
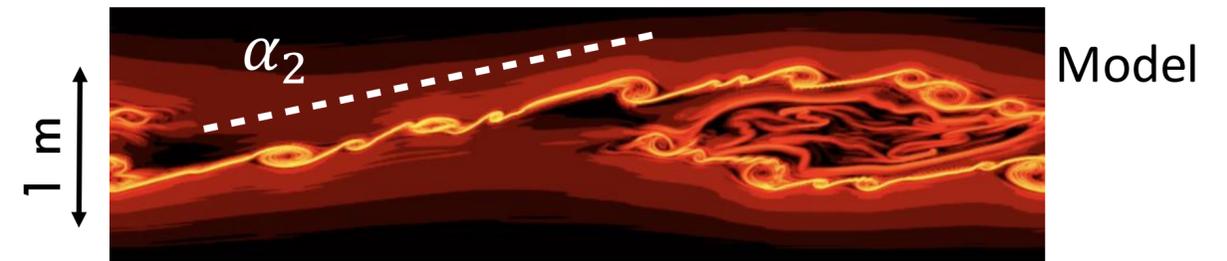
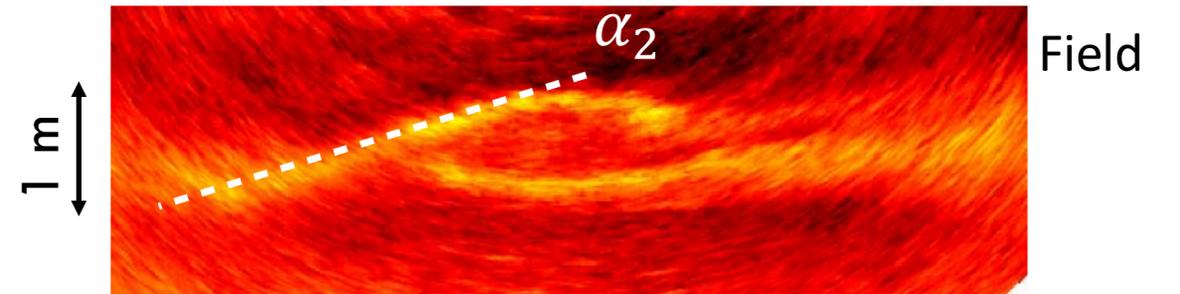
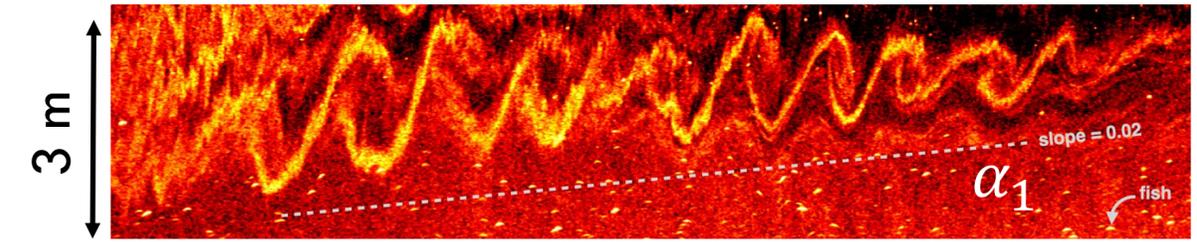


# Modelling the braid mixing rate and its small-scale structure



# Takeaways

1. Shear instability under continuous shear forcing ( $\mathbf{Ri} \approx 0.1 - 0.15$ ) mix along thin, sloping braids rather than through large overturns
2. Multibeam echo-sounding disentangles spatial and temporal evolution, confirming shallow braid slopes  $\alpha \approx 0.2 - 0.4$
3. At high  $Re$ , stratified braid baroclinicity triggers fast secondary shear instabilities and turbulence that pre-empt primary overturn
4. A simple baroclinic production–dissipation balance predicts  $\epsilon$  and  $\chi_s$  at both pycnocline and braid scales as a function of local  $\alpha, \delta, S, N$
5. Elevated braid  $\epsilon$  and  $\chi_s$  predictions match the field, and lab data reveal extreme  $\chi_s$  in fine sub-mm filaments, not overturns

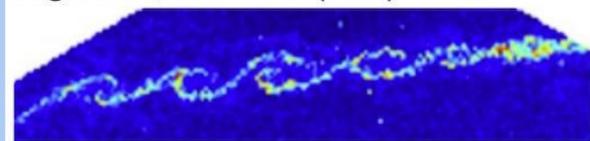


Find out more: preprint



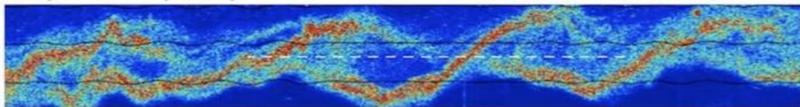
# Gallery of shear instabilities

**Knight Inlet, Colbo *et al.* (2014)**

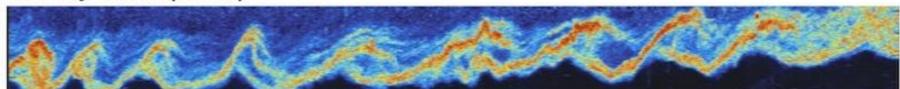


**Connecticut River estuary,**

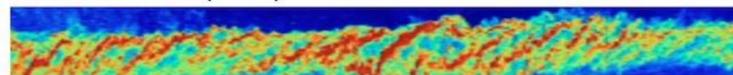
(a) Geyer *et al.* (2010)



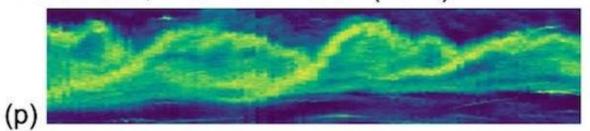
(b) Lavery *et al.* (2013)



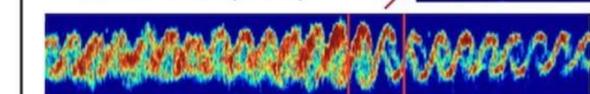
(c) Holleman *et al.* (2016)



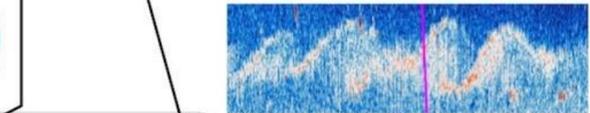
**Baltic Sea, Muchowski *et al.* (2022)**



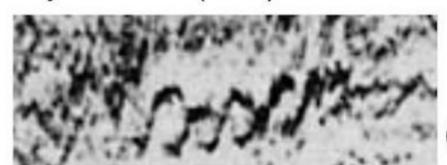
**Changjiang River estuary,**  
(q) Tu *et al.* (2022)



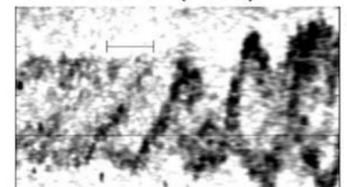
(r) Tu *et al.* (2024)



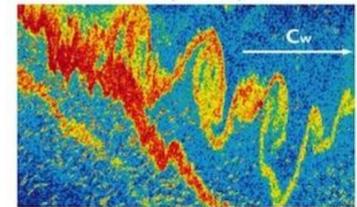
**Fraser River estuary,**  
Geyer & Smith (1987)



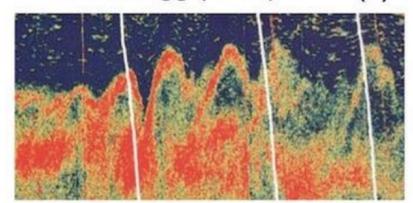
Tedford *et al.* (2009)



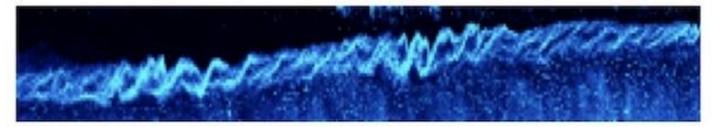
**Oregon Shelf,**  
Moum *et al.* (2003)



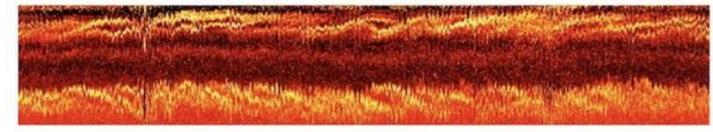
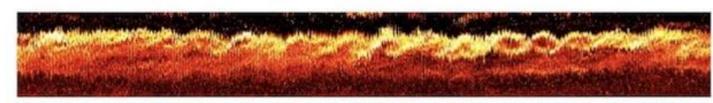
**Admiralty Inlet,**  
Seim & Gregg (1994)



**Columbia River estuary,** Lavery (2023) (i)



**Mobile Bay,** Bassett (2021)



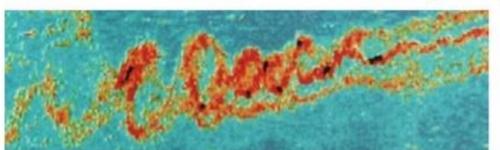
**James River Estuary,** Bassett *et al.* (2023)



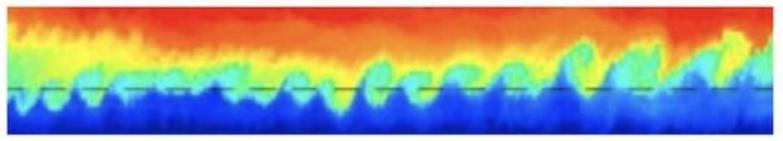
**New England Shelf,**  
Zhang & Lavery (2016) (o)



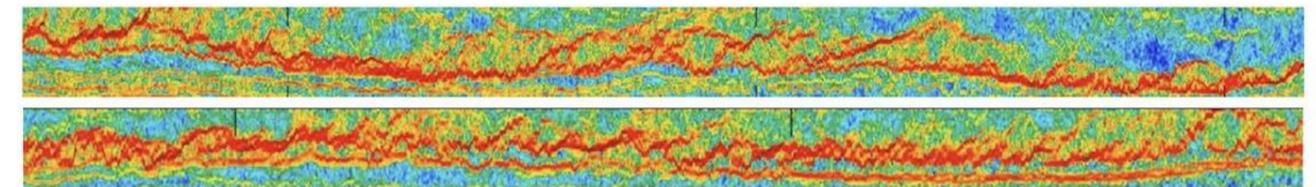
(n) **Strait of Gibraltar,** Wesson *et al.* (1994)



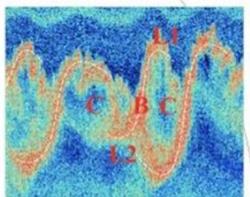
(m) **Great Meteor Seamount,**  
van Haren & Gostiaux (2010)



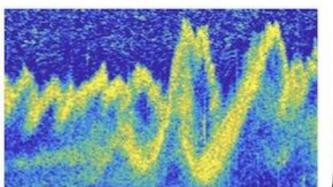
(l) **Romanche Trench,** van Haren *et al.* (2014)



**Taiwan Seamount,**  
(s) Chang *et al.* (2016)



(t) Vladioiu *et al.* (2025)



(u) **Samoa Passage**  
Cusack *et al.* (2019)

